

Utilisation of inorganic and organic soil phosphorus in a hill country soil

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Abstract

Changes in phosphorus (P) fractions of unfertilised and fertilised (superphosphate) soil were investigated over five years at a hill country site near Te Kuiti. Only soil inorganic P (Pi) reserves were utilised for plant uptake when superphosphate was withheld at the site. Immobilisation of P as soil organic P (Po) contributed to depletion of the soil Pi reserves during the first two years of this trial. Where superphosphate was applied, immobilisation of P as Po amounted to about 25% of applied P during the five years measurements were made. Changes in soil P fractions indicated that all forms of soil Pi were utilised when superphosphate was withheld. These included readily available Pi, Al-Pi, Fe-Pi, and residual phosphate rock from previous fertiliser applications. Depletion of the phosphate rock residues in the soil also occurred where superphosphate was applied and appears to have been completed within about two years. The phosphate rock residues had probably accumulated because of the relatively high amounts of unacidulated phosphate rock in superphosphate manufactured before 1983. Accumulation of Po associated with humic acid, or adsorbed on surfaces of hydrous oxides of Al and Fe, occurred in both fertilised and unfertilised soils. The more labile forms of Po also increased in the fertilised soil.

Keywords inorganic phosphorus, organic phosphorus, phosphorus immobilisation, soil phosphorus, soil phosphorus fractions, soil phosphorus utilisation.

Introduction

Fertiliser use by New Zealand agriculture in the late 1980s was 50 to 60% below that in the late 1970s (Taylor *et al.* 1989). The consequent decline in soil fertility was particularly evident in hill country farms. Significant decreases in animal and pasture production occurred on hill country pastures where fertiliser was withheld (Gillingham *et al.* 1990; Lambert *et al.* 1990; O'Connor *et al.* 1990). In this situation, nutrient reserves accumulated from past fertiliser applications determined soil fertility and pasture production.

Fertiliser application (mostly superphosphate), together with biological fixation of nitrogen by legumes, has encouraged the buildup of organic matter in New Zealand pastoral soils. Phosphorus (P) has accumulated in both inorganic (Pi) and organic (Po) forms (Perrott *et al.* 1989). Po amounts to 30-80% of total P in New Zealand pastoral soils (Perrott and Sarathchandra 1987).

As these P reserves determine soil P status when fertiliser is withheld, the extent to which the different P forms are utilised is important. This paper reports the effects of plus and minus fertiliser treatments on changes in soil P fractions and the utilisation of soil Pi and Po at a hill country site near Te Kuiti.

Materials and methods

A small plot trial was established on two paddocks that were part of the nil fertiliser treatment in an animal grazing experiment (O'Connor *et al.* 1990) at the AgResearch, Te Kuiti Research Area, 20 km south of Te Kuiti.

The soil was a Mahoenui silt loam, a central yellow-brown earth (Typic Dystrachrept) typical of 1.3 million ha of North Island hill country. A thin veneer of volcanic ash was present in the surface layer of this soil. Samples (0-75 mm) taken at the first sampling in August 1983 had a med-high Pretention (68%) and a mean Olsen P of 14. The 0-30 mm and 0-75 mm depths had average bulk densities of 0.54 and 0.66 g/cm³ respectively. Superphosphate had been applied (average rate 250 kg/ha/year) for at least 10 years before the animal grazing experiment which began in April 1983. The pasture was grazed by a flock of sheep (stocking rate 14 ewes/ha) restricted to these and four other paddocks that received no fertiliser.

Six control and six fertilised (250 kg superphosphate/ha/year) paired plots were established in each paddock. Superphosphate was applied to the fertilised plots immediately after the first soil sampling in August 1984 and in September 1985, 1986, 1987, and 1988. Pasture production was determined for each plot by a trim cage technique and herbage samples from each cut were analysed for P.

Long term fertiliser effects were investigated with soil samples taken to 30 mm depth (15 cores 25 mm dia) at approximately three monthly intervals from August 1984 to June 1989. Samples were also taken to 75 mm depth at 4- to 7-weekly periods from August 1984 to

December 1986 as part of a separate study (Perrott *et al.* 1992) and these were analysed for Pi and Po. After air drying the soil samples were sieved (<2 mm) to remove roots and plant debris. They were then further ground and sieved (<150 μm).

With the 0-75 mm samples, Pi was determined by extraction with 0.5 M H_2SO_4 (almost equal to total Pi in this soil) and total Po by ignition (Saunders and Williams 1955; Blakemore *et al.* 1972). A sequential extraction procedure (Perrott *et al.* 1989) was used to fractionate Pi and Po in the 0-30 mm samples. This involved extraction of soil Pi and Po with 0.5 M NaHCO_3 (pH=8.5), followed by 0.1 M NaOH and then 1.0 M NaOH . Remaining Pi was extracted with 0.5 M H_2SO_4 . Residual Po was then determined after ignition (550°C) and extraction with 0.5 M H_2SO_4 . To prevent precipitation of Ca-Pi and Ca-Po compounds in the alkaline extracts (Perrott 1992), exchangeable Ca was removed from the soil samples before the sequential extractions by washing with 1 M NaCl .

Data analysis

Differences between fertilised and unfertilised plots were tested by analysis of variance at each sampling date. Rates of change of total P and the different P fractions for each plot were calculated using regression analysis against time (August 1984 - August 1986 for 0-75 mm samples and August 1984 - June 1989 for 0-30 mm samples). Means and standard errors of means (SEM) were then calculated for unfertilised and fertilised soil.

Results and discussion

Pasture production and P uptake

Withholding fertiliser reduced pasture production slightly in the first year of the trial, but there were reductions of approximately 20% in subsequent years (Table 1). The decline in P uptake was more marked, ranging from approximately 20% reduction in the first year to approximately 40% in the fifth year. Pasture composition was also influenced by withholding fertiliser, with a decline in white clover content and an increase in moss and weeds (O'Connor *et al.* 1990).

Total P, total Pi and total Po

The changes in soil Pi and Po with time (0-30 mm, 1984-1989) are presented in Figs. 1 and 2. Table 2 summarises the original levels of Pi, Po and total P, and their rates of change with time, for both time periods using units of kg P/ha in the 0-75 mm depth. Data for the 1984-1989 period were estimated from the 0-30 mm samples using

Table 1 Effect of withholding fertiliser on pasture production and P uptake

Year	Unfertilised	Fertilised	SED	%decline
Pasture production (kg DM/ha/year)				
84/85	6220	6630	343	6
85/86	7300	9100	445	20
86/87	7460	9960	450	25
87/88	8720	10710	497	19
88/89	7010	8700	351	19
P uptake (kg P/ha/year)				
84/85	13.6	17.2	0.89	21
85/86	17.2	25.0	1.36	31
86/87	14.4	23.9	1.08	40
87/88	20.2	30.3	1.29	33
88/89	14.6	25.8	1.10	43

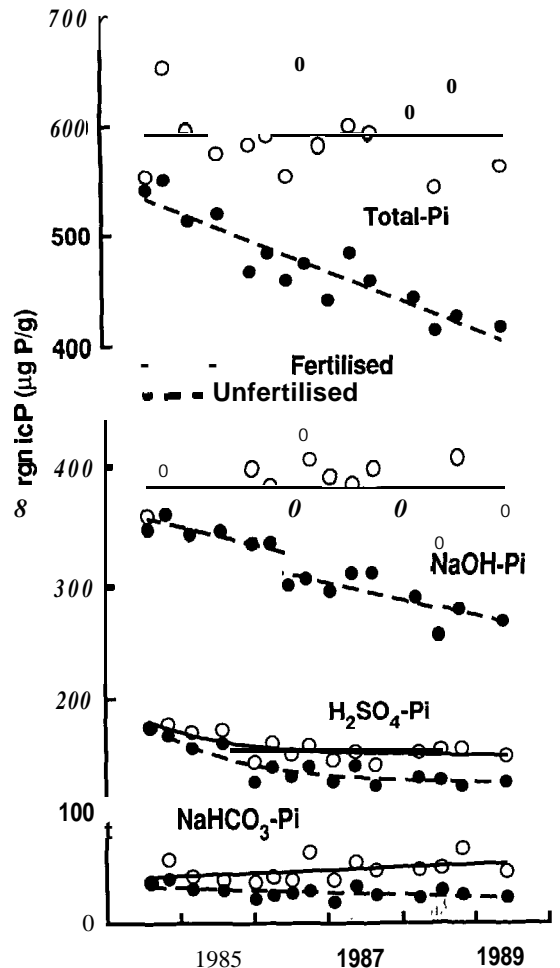


Figure 1 Changes in total soil pi and Pi fractions with time (0-30 mm samples). closed symbols - unfertilised plots. Open symbols - fertilised plots. Lines are linear fits to data except for H_2SO_4 -Pi where an exponential fit was used.

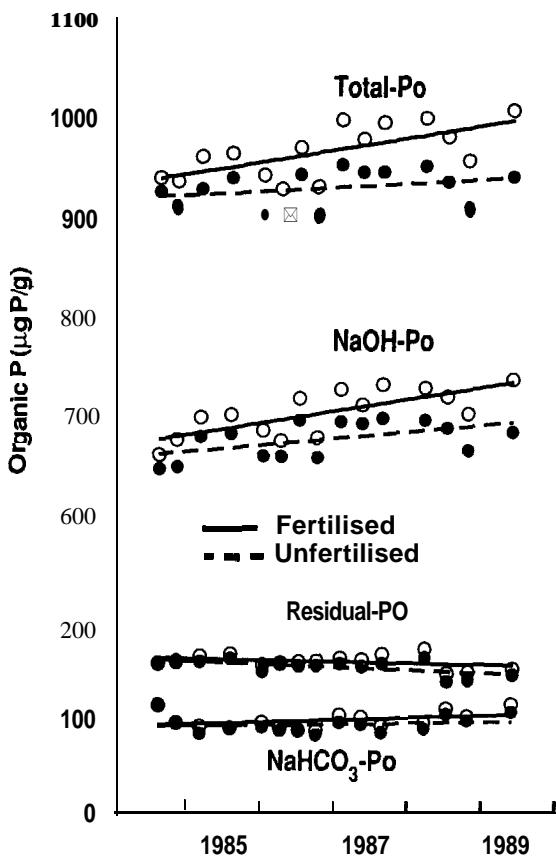


Figure 2. Changes in total soil Po and P fractions with time (0-30 mm samples). See Fig.1 for explanation of symbols. Lines are linear fits to data.

Table 2 Estimated rates of change in soil P levels (kg P/ha/year for 0-75 mm depth)

	Original level	1964-1969 estimated from 0-30 mm data ^a (n=15)		1964-1966 estimated from 0-75 mm data (n=19)	
		Unfertilised	Fertilised	Unfertilised	Fertilised
Total Pi	190	-9.7	-0.5	-12.9	-9.1
(SEM)		(0.67)	(1.21)	(1.8)	(2.3)
Total PO	480	+2.0	+1.8	+12.1	+10.0
(SEM)		(1.06)	(0.88)	(1.7)	(3.2)
Total P	670	-10.5	+4.6	-0.6	+1.0
(SEM)		(1.16)	(2.1)	(2.23)	(5.07)

a. Estimated from 0-30 mm data using mean ratios of Po, Pi and total P in 0-75 mm samples to total extractable Po, Pi and total P in 0-30 mm samples for eight sampling dates in 1964-1966.

the mean ratios of Po, Pi and total P in the 0-75 mm samples to the total extractable Po, Pi and total P in the corresponding 0-30 mm samples (eight sampling dates 1984-1986). This was done because the 0-75 mm depth is more representative of the rooting zone than is the 0-30 mm depth.

There were no significant trends in total P level of samples taken from either the fertilised or unfertilised plots during 1984-1986 (0-75 mm). However, Pi decreased and Po increased during this period for both the fertilised and unfertilised plots. During 1984-1989 the total P and total Pi levels (0-30 mm) of the unfertilised plots declined, while total P increased in the fertilised soil. There were no significant trends for total Po in the unfertilised plots or total Pi in the fertilised plots over the 1984-1989 period.

These changes in soil P levels indicate that Pi reserves were the main source of P for plant uptake when fertiliser was withheld at this site. There was no net utilisation of soil Po. The decline of soil Pi by about 10 kg P/ha/year is approximately 5% of the total Pi reserves of 190 kg P/ha (0-75 mm). Immobilisation of fertiliser P as soil Po in the fertilised plots (about 6-10 kg P/ha/year, 0-75 mm) was a significant component of fertiliser P application (25 kg P/ha/year). Without fertiliser, significant immobilisation of P as Po occurred during the first two years. This would have contributed to the depletion of soil Pi reserves.

Net stock transfer of P away from the plots (the rest of the paddock was not fertilised), and immobilisation of P as soil Po, can explain the failure of Pi to increase in the fertilised soil. Because the fertilised plots were inside unfertilised paddocks all the net pasture uptake of P due to fertiliser would have been transferred by stock. This amounted to about 35% of applied P (Table 1) and together with the net changes in total soil P (Table 2) accounted for all the fertiliser P applied. In a situation where the whole paddock was topdressed, net stock transfer would be smaller and soil Pi levels would probably increase.

The relatively low level of soil Pi reserves explain the marked decline in pasture production and P uptake at this site. In a similar study Perrott et al. (1990) measured soil Pi reserves of about 960 kg P/ha (0-75 mm) at a high-producing dairy farm on a yellow-brown loam in the Waikato. Pasture production at this site did not decline until seven years after stopping fertiliser when a 8% decrease was measured. The rate of soil Pi decline at that time was approximately 24 kg P/ha/year (0-75 mm).

Soil Po levels at the Waikato site were approximately 850 kg P/ha (O-75 mm) which was much higher than for the Te Kuiti soil. No significant changes in soil Po with time were observed at the Waikato site, indicating that the soil was at, or near, equilibrium with respect to **Po** accumulation. In contrast, Po was still accumulating at the Te Kuiti site. Nguyen *et al.* (1989) also demonstrated accumulation of Po in fertilised **and unfertilised** soil at an irrigated mid-Canterbury site which had similar Po levels to the Te Kuiti soil. The rate of accumulation of soil Po (5.8 kg P/ha/year) was similar. Apparently, net accumulation of soil Po under pasture is likely until an equilibrium level is attained as has evidently occurred with the Waikato soil. Given the relatively low levels of fertiliser topdressing on hill country farms, it is likely that most hill country pastoral soils are still at a stage where net immobilisation of P as Po is occurring.

Pi and Po fractions

Phosphate fertiliser effects concentrate in the surface layers of the soil. A 0-30 mm sampling depth was therefore chosen to reduce dilution of any fertiliser effects on the soil P fractions. The sequential extraction procedure used (Perrott *et al.* 1989) was designed to separate soil P into pools of decreasing availability to plants (Hedley *et al.* 1982). The initial, milder reagents extracted the more labile forms of soil Pi and Po.

Table 3: Rates of change of soil Pi and Po fractions with time ($\mu\text{g/g/year}$ for 0-30 mm depth)

	Unfertilised		Fertilised	
	Pi	Po	Pi	Po
NaHCO₃	-2.1	+0.6	+2.4	+1.9
(SEM)	(0.35)	(0.73)	(0.81)	(0.62)
NaOH	-14.9	+6.3	+1.5	+11.5
(SEM)	(1.25)	(1.41)	(2.30)	(1.30)
H₂SO₄-Pi	See note A		See note A	
Residual-Po		See note B		See note B
Total	-26.4	+3.8	-1.4	+11.8
(SEM)	(1.62)	(2.06)	(3.56)	(1.72)

A. See Fig. 1. A significant decline occurs over first two years.
 B. See Fig. 2. No significant change occurs in first four years.

Amounts of each soil P fraction are presented for each sampling time in Figs. 1 and 2. The 0.1 M NaOH and 1.0 M NaOH fractions have been combined. Table 3 summarises trends in the change of soil P fractions with time.

In the unfertilised soil, all the Pi fractions declined significantly with time (by about 4-6% per year) indicating utilisation of all forms of soil Pi readily

available (**NaHCO₃-Pi**), less available forms of Al-Pi **and** Fe-Pi (**NaOH-Pi**) and apatite P (**H₂SO₄-Pi**). However, most of the decline in **H₂SO₄-Pi** occurred in the **first** two years (Fig. 1). This probably represents dissolution of unacidulated phosphate rock accumulated from previous superphosphate topdressing.

A similar net dissolution of accumulated phosphate rock residues also appeared to occur in the fertilised soil despite addition of some unacidulated phosphate rock in **current** superphosphate applications. This probably **resulted** from higher **than** normal amounts of unacidulated phosphate rock in New Zealand superphosphate manufactured between 1970 and 1983 (Edmeades *et al.* 1990). Increasing amounts of **NaHCO₃-Pi** reflected **superphosphate** addition but there was no significant trend for **NaOH-Pi**. Annual application of soluble P, which is adsorbed by hydrous Al and Fe oxides, increased the variability of the **NaHCO₃-Pi** and **NaOH-Pi** measurements.

The last three sampling dates strongly effected the regression analysis for residual Po which included the most recalcitrant forms of soil Po. However, no **significant** trends occurred over the **first** four years (Fig. 2). Po associated with **humic** acid or adsorbed by hydrous oxides of Al and Fe (**NaOH-Po**) increased significantly with time in both the unfertilised and fertilised soils. There was also a smaller, but significant, increase of the more labile soil Po (**NaHCO₃-Po**) in the fertilised soil.

While there was no net utilisation of soil **Po** at this site, mineralisation of the more labile **Po** compounds could still have occurred. Input of labile **Po** compounds from plant roots and soil micro-organisms may have replaced losses due to mineralisation. Accumulation of soil Po probably occurs as compounds like phytates that are not mineralised when adsorbed by soil surfaces.

Conclusions

Withholding fertiliser at this site caused depletion of the soil Pi reserves by about 10 kg P/ha/year (O-75 mm) equivalent to 5% of total Pi reserves at this depth. The depleted reserves included readily available Pi, Pi associated with soil Al and Fe, and unacidulated phosphate rock accumulated from past superphosphate applications. No net depletion of soil **Po** occurred. In the fertilised soil approximately 25% of applied P was immobilised as soil **Po**. Accumulation of soil **Po** also occurred with unfertilised soils during the **first** two years of the trial.

The low Pi reserves and net accumulation of **Po** is probably common to most hill country pastoral soils because of past fertiliser history. Consequently, continued inputs of fertiliser P are required on these soils to maintain existing Pi reserves and prevent production declines.

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